Core formation in planetesimals: the importance and geochemical legacy of immiscibility in metallic liquids.



Geoffrey Bromiley (geoffrey.bromiley@ed.ac.uk), Freya Dibble, Lee Saper School of GeoSciences, The University of Edinburgh, UK.

Segregation of Fe-rich liquids and formation of metallic cores was a key differentiation process in rocky bodies in the early solar system. Mechanisms for

-The dominant process of core formation was probably segregation of Fe-rich liquid from a crystal-rich magma ocean (left).

segregation in planetesimals (100-1000km bodies), and the proportion of these

bodies which differentiated, remain unclear. Here we explore the important role

of immiscibility of S-rich and S-poor liquids in core formation models.

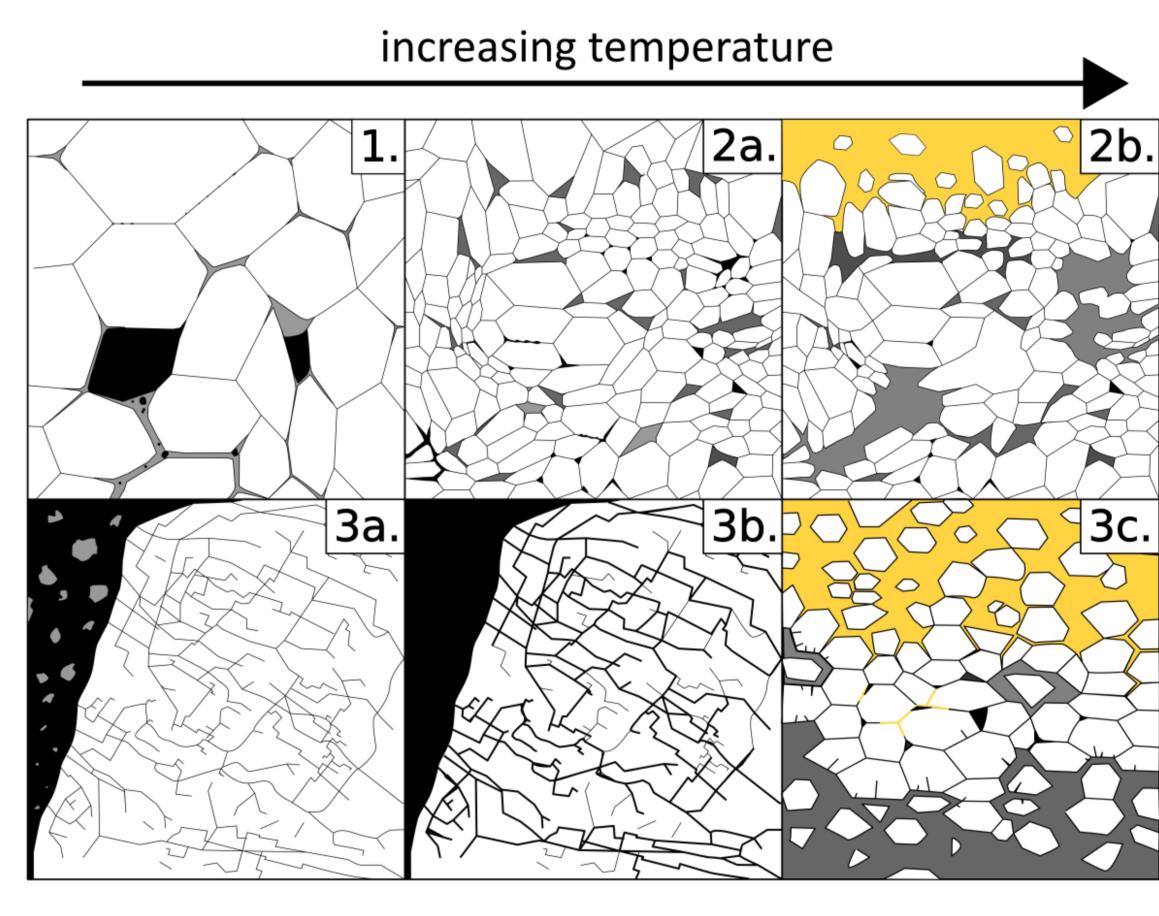
- -In S- and P-rich systems¹ and at C saturation^{2,3} immiscibility results in separation of S-rich and S-poor core-forming liquids.
- -Various 'low-temperature' models of core-formation have also been proposed (below right) which typically involve percolation of FeS-rich liquids through crystalline silicate, either before or just after the onset of low-degree silicate melting.
- -Mechanisms and conditions of segregation will have strong influence on core composition; e.g. early-formed (low T) metallic liquids are assumed to be S-rich (near eutectic in the Fe-FeS system), with a decrease in S content at higher T. However, immiscibility can also result in segregation of S-rich liquids.

Fe-rich liquid
FeS-rich liquid
Silicate magma ('mush')
Silicate-rich, crystalline

For key see references below

Low-temperature core-formation models:

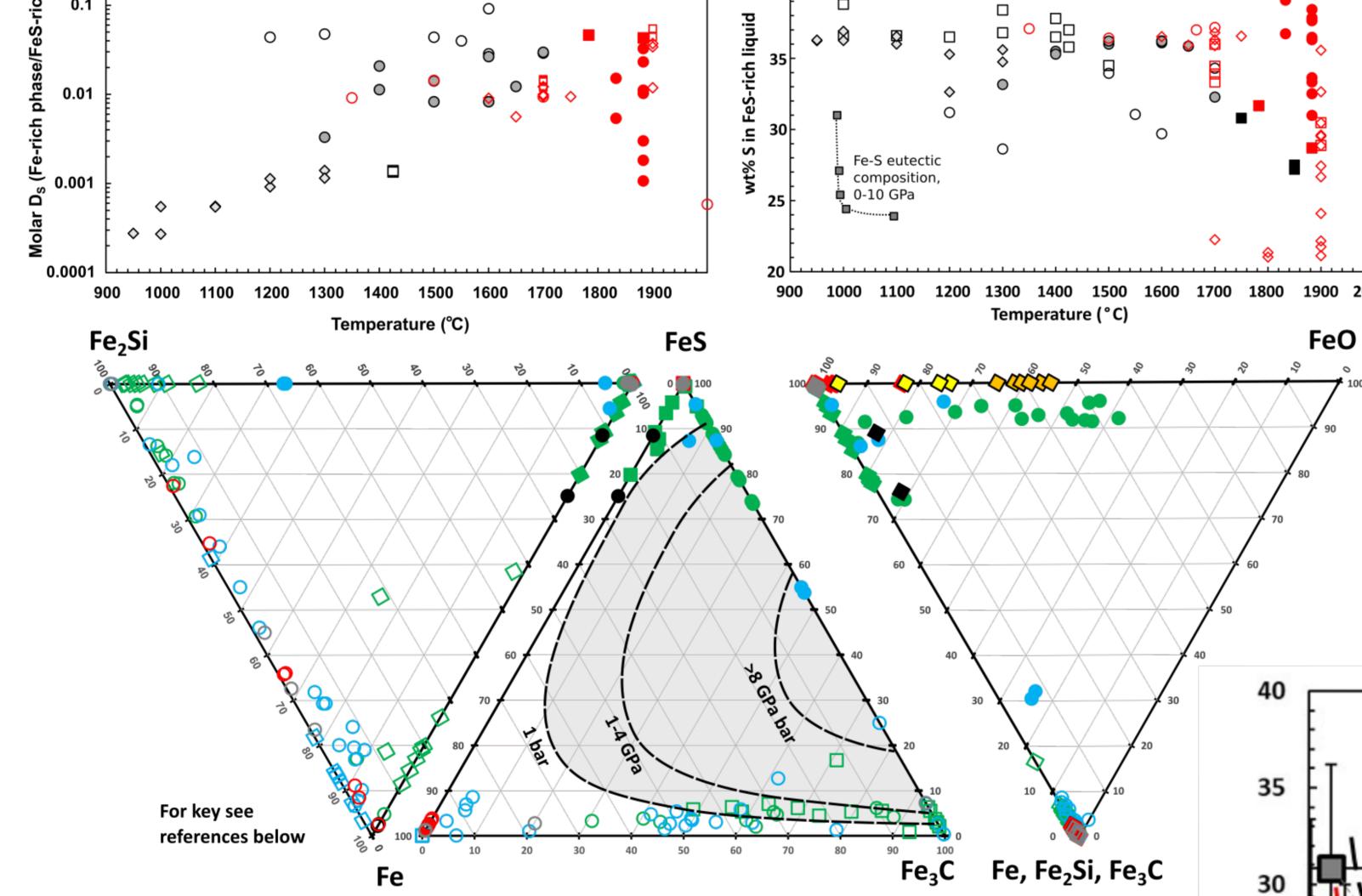
- 1. Percolation of S-rich liquids, aided by deformation, porosity etc.
- 2. Low-degree silicate melting increases metallic melt fraction and aids phase segregation.
- 3. Tomkins et al.⁴ model of impact driven FeS liquid mobilisation and subsequent segregation at the onset of silicate melting.



A review of published data from partial melting experiments in chondritic systems (left) provides insight into the composition of metallic liquids formed under various conditions:

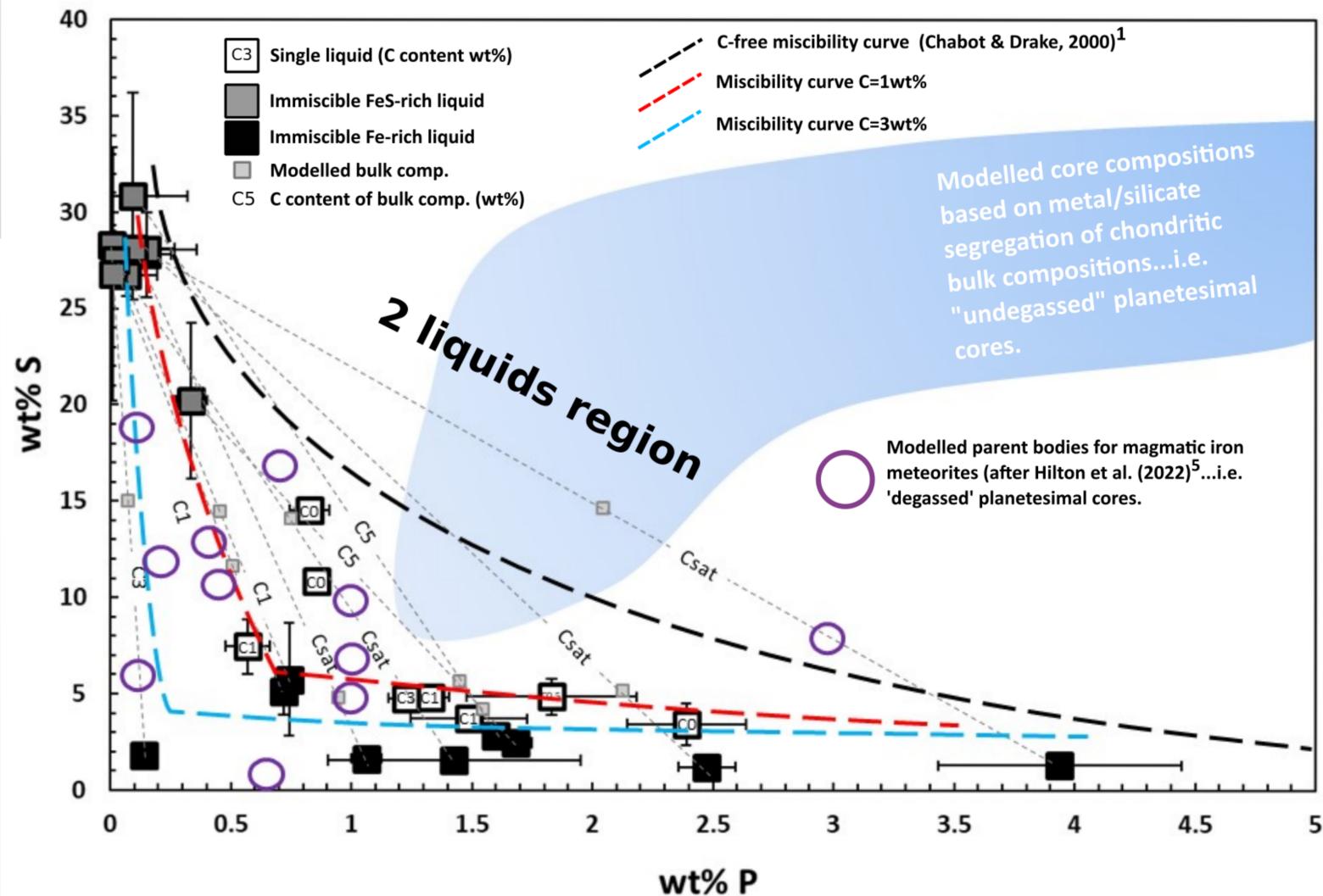
- (1) as expected, initial liquids are S-rich, with a large T range over which metallic components melt before the onset of silicate melting (left, top). However,
- (2) FeS-rich liquids have S contents exceeding Fe-FeS eutectic compositions, with little change in composition with increasing T, and
- (3) There is immiscibility/separation of Fe-rich (S-poor, P-rich, C-rich, O-poor, Si-bearing) and FeS-rich (P-poor, C-poor, O-bearing) solids/liquids (left, bottom) over a wide P-T-fO₂ range.

Partial melting experiments are likely C-saturated due to experimental design, meaning that metallic liquid compositions and melt relations are dominated by the effects of immiscibility.



-Our new superliquidus experiments (0.5 GPa/1673K) in the system Fe-Ni-P-S-O-C **constrain the role of C in driving immiscibility (right).** Bulk C contents resulting in immiscibility mirror trends in C solubility in Fe-rich liquids⁶, although immiscibility occurs at C contents considerably below saturation.

- -Differentiation of planetesimals with chondritic compositions (assuming no degassing/loss of C, P, S) results in core immiscibility.
- -In contrast, parent bodies to magmatic iron meteorites were single liquids due to substantial degassing⁷ and low C, P, S contents.
- -Bulk composition, size and growth rate, and degree of degassing, controlled the extent of immiscibility in planetesimal cores.
- -Immiscibility can also occur during low-temperature core formation, and obfuscates results from partial melting experiments.



References: ¹Chabot NL & Drake (2000) Meteorit. & Planet. Sci. 35:807. ²Corgne et al. (2008) GCA 72:2409. ³Dasgupta et al. (2009) GCA 73:6678. ⁴Tomkins et al. GCA 100:41. ⁵Hilton et al. (2022) GCA 318:112. ⁶Zhang et al. (2018). GCA 225:66. ⁷Grewal DS & Asimow GCA 344, 146–159 (2023). Black/grey =data from partial melting experiments. Open squares: McCoy et al. (1999) Meteorit. & Planet. Sci. 34:735, circles: Berthet et al. (2009) GCA 73:6402, diamonds: Ford et al. (2008) Meteorit. & Planet. Sci. 43:1399, black squares: Corgne et al. ²Red =data from FeS/silicate liquid partitioning experiments. Red squares: Steenstra et al. (2020) GCA 286:248, unfilled squares: Boujibar et al. (2019) Am. Min. 104:1224, open circles: Malavergne et al. (2014), red circles: Steenstra et al. (2020) Icarus 335, open diamonds: Boujibar et al. (2020) GCA 269:622. Ternary plots. Blue squares: McCoy et al. (1999), green squares: Berthet et al. (2009), red squares: Ford et al. (2008), green circles: Boujibar et al. (2020), blue circles: Boujibar et al. (2019), black squares: Steenstra et al. (2020a), red circles: Steenstra et al. (2020b), grey circles: Malavergne et al. (2014) EPSL 394:186, yellow squares: FeS from percolation experiments of Terasaki et al. (2008) EPSL 273:132. Shaded region on FeS-Fe-Fe₃C ternary: immiscibilities from Corgne et al. (2008)².